The modeling way-back machine

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Remember what it was like back them! This is what my computer looked like:









I was at NASA Ames...







And, we all knew you couldn't scale from leaf to canopy!





Partly, because this guy told us so...

State of the art ca 1985: Growth curves simulating annual biomass accumulation and NPP, no fluxes



There were at least two big reasons:

- Canopies were t complex to model without scaling rules (brute force didn't work very well, the data to describe a canopy was laborious to collect and didn't generalize).
- 2. We didn't know the answer (we could model fluxes but there was no way to tell if the simulation got the right answer)



These were heady times for modeling, though, the key parameterizations we still use were new then:

Growing plants



$$\begin{split} V_{c} &= V_{c_{\max}} \frac{C}{C + K_{c}(1 + O/K_{o})} \cdot \frac{R}{R + K_{r}'} \\ &= V_{c_{\max}} \frac{C/K_{c}}{1 + C/K_{c} + O/K_{o}} \cdot \frac{R/K_{r}'}{1 + R/K_{r}'} \end{split}$$

Harvest SYMBIOTIC N HARVEST REMOVAL FIXATION ABOVE-GROUND BELOW-GROUND ERMVST(N) SNFXAC EGRAININ IVE HERBACEOUS LIVE HERBACEOUS VOLATILIZATION N AGLIVE(N) BGLIVE(N VOI PI Tillage STANDING DEAD RI FAVEIN FROOTE(N HERBACEOUS N Fertilizer N STDEDE(N) H Death Tillage 📯 FERTOT(N) C/N DA Organic N Fall Rate 🕅 Tillage Addition Volatilization OMADAE(N) Death N.O. NO., Plant Residue Plant Residue Atmospheri C/N Lignin to C/N C/N Lignin to C/N Deposition WDFXA SURFACE SURFACE SOIL SOIL STRUCTURAL N METABOLIC N METABOLIC N STRUCTURAL N STRUCE(2, N) STRUCE(1, N) METABE(1, N) METABE(2, N) Non-Symbiotic MINERAL N N Fixation MINERL(1.N min Rad M X C/N Sand Leaching CTIVE SURFACE ACTIVE SOIL DEAD FINE MICROBE N MA ORGANIC N SOM1E(N) BRANCH N SOM1E(2, N NLAYER · WOOD1E(N NI AVER Rad ∽ N_{min} ŴC/N MINERL C/N 🕅 M (LAYER, N) Lignin STREAM(2) DEAD LARGE C/N, M, WOOD N SLOW SURFACE WOOD2E/M SLOW SOIL Mixing ORGANIC N ORGANIC N SOM2E(1, N) SOM2E(2, N) C/N DEAD COARSE Lignin C/N. M ROOT N Sand WOOD3E(N) C/N Clay Leaching A Compost M = multiplier for effects of moisture, temperature, cultivation, pl PASSIVE SOIL Drganic N Addition OMADAE(N) I FACHED Nmin = N mineralization from CO2 loss ORGANIC N ORGANIC N Rad = solar radiation STREAM(6) SOM3E(N)

Decomposing plants

Parton soil organic matter

Farquhar-Berry photosynthesis

Seminal... an overused word but right on here!

<u>A biochemical model of photosynthetic CO2 assimilation in leaves of C3 species</u> <u>GD Farquhar, S von Caemmerer, JA Berry</u> - Planta, 1980 - Springer Abstract Various aspects of the biochemistry of photosynthetic carbon assimilation in C 3 plants are integrated into a form compatible with studies of gas exchange in leaves. These aspects include the kinetic properties of ribulose bisphosphate carboxylase-oxygenase; ...

5494 citations

Analysis of factors controlling soil organic matter levels in Great Plains grasslands

WJ Parton, <u>DS Schimel</u>, CV Cole... - Soil Science Society of ..., 1987 - dl.sciencesocieties.org

Abstract We analyzed climatic and textural controls of soil organic C and N for soils of the US Great Plains. We used a model of soil organic matter (SOM) quantity and composition to simulate steady-state organic matter levels for 24 grassland locations in the Great Plains. ...

3320 Citations



Empirical evaluation at the Konza and environs! Mark







Tim



Using the landscape



FIG. 1. An idealized cross section through a Konza watershed. Soil depth increases downslope, but incision to bedrock may occur in the ephemeral channel. Loess caps may occur on ridgetops as shown, although not all ridgetops have such deposits. Transects used in this study spanned four such toposequences; physiological studies were carried out in lowland positions, usually just above the drainage channels, and on steep limestone sites at slope shoulders.

I'm pretty sure this picture was done entirely without benefit of a computer



Data across landscapes



FIG. 3. Biomass and leaf area index (LAI) data from transects shown in Fig. 2. Surfaces are plotted as a function of transect distance and time (the four IFCs in 1987). Sample points and IFC dates are indicated by numbered posts. (A) and (B) Live aboveground biomass for burned and unburned transects, respectively. (C) and (D) LAI across the same transects. LAI was computed using equations from Table 1.



Nitrogen (V_{cmax}) across landscapes.



FIG. 4. N data from transects shown in Fig. 2, plotted against transect distance and time as in Fig. 3. (A) and (B) N mass in live vegetation, per unit sample area. (C) and (D) N per unit leaf area (N mass per unit sample area LAI), a correlate of maximum photosynthetic rate. There was no live biomass at IFC 4. Note that (C) and (D) were rotated to best display dynamics.



The relationship between light interception and nitrogen limitation



FIG. 6. The N gradient index ([N] in the top canopy layer/ [N] in the bottom canopy layer) plotted against total IPAR (expressed as fraction of incoming PAR) for 1988 and 1989. As IPAR increased, the gradient in N allocation within the canopy became steeper.



V_{cmax} scaling in real canopies



FIG. 9. Photosynthetic capacity by height for unburned (hatched) and burned (stippled) canopies. Note the rapid decline in capacity with height and the steeper gradient of photosynthetic capacity (top: bottom = 3.19) in the taller, more productive unburned sites compared to the unproductive burned sites (1.53). Error bars show 1 sp. Data are from 1989.



Light response curves, stratified by depth in the canopy





Scaling rules: light extinction and photosynthetic capacity Simulated response curves

the general pattern matches the prediction from theory $-A_{max}$ and leaf nitrogen [N] - should scale with the time-integral of the absorbed local PAR



or

Scaling rule, \checkmark , now, what's the answer?

- Knowing how A scales through a canopy, can we get the right daily, weekly, annual GPP?
- It remains complex to integrate over diurnal, synoptic and seasonal time scales..
- FIFE provided the first answers, sustained eddy covariance data!



Pioneering eddy covariance research in FIFE..

A few reminders:

- FIFE Eddy covariance data provided the first flux data from a major integrated campaign and over ecologically meaningful time scales.
- Flux data from FIFE have been used by nearly every modeling group to develop improved A – R and ET models.





Still in routine, daily use by modelers...



A case study for land model evaluation: Simulation of soil moisture amplitude damping and phase shift (Wu, Geller and Dickinson, 2002)



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http://onlinelibrary.wiley.com/doi/10.1029/2001JD001405/full#jgrd9437-fig-0001

Work at the Konza established linkages between the biophysical and biogeochemical that are <u>still</u> being fully absorbed into models

- We still don't fully understand nutrient limitation,
- Links between nutrient and physical (light, water) controls remain a frontier
- Does nutrient limitation equilibrate to energy availability (Schimel et al 1997) or does the geochemical evolution of soil control energy harvesting?

So, scaling from leaf to canopy, routine, and now the standard for modeling, we're now combining leaf/canopy and individual-based modeling

Ecosystem Demography Model (ED2)



Traditional 'big leaf' model v.s. Age and Size structured model

As we enter the era of forest demographic models to address longer timescales and changing disturbance regimes

Lessons learned

- FIFE catalyzed the fusion between atmospheric and ecosystem science via radiation and turbulence (a hallmark of NASA programs): made a huge contribution to Earth System Science.
- FIFE resulted in two breakthroughs in canopy scaling theory, and canopy model validation that enabled both land surface and remote sensing-driven modeling for decades to come.
- FIFE established a precedent of close partnership between modelers and observationalists that has accelerated the pace of science.
- There are still open questions hidden in plain site in FIFE, BOREAS and LBA data!
- Field Campaigns are addictively fun-many of us are still doing them long after we should know better.