

Atmospheric and Environmental Research



## OSSE Towards Quantifying ASCENDS Measurement Requirements to Constrain North American Regional CO<sub>2</sub> Fluxes

Janusz Eluszkiewicz, Marikate Mountain, Ryan Aschbrenner, John Henderson Thomas Nehrkorn, Jennifer Hegarty, and Scott Zaccheo

Atmospheric and Environmental Research, Inc.

February 28, 2012

# **Overview of Regional Method**

•Apply the <u>S</u>tochastic <u>Time-Inverted Lagrangian Transport</u> (STILT) model driven by meteorological fields from the <u>W</u>eather <u>R</u>esearch and <u>F</u>orecasting (WRF) model to generate surface influence functions, "footprints", for ASCENDS observations.

•The "footprints" (or adjoint) express the sensitivity of ASCENDS column  $CO_2$  observations to surface fluxes in the upwind source regions.

•Footprints enable the computation of posterior flux error reductions resulting from the inclusion of ASCENDS observations.



# **ASCENDS Candidate CO<sub>2</sub> Wavelengths**

### 3 Concepts:



### **Vertical Sensitivity**

Atmospheric and Environmental Research

# **ASCENDS Candidate CO<sub>2</sub> Wavelengths**

### Measurement Errors for 1.57µm vs 2.05µm

1 - 31 Jan -- 1.57 μm Line Min: 3.26 ppmv \_\_\_\_ Max: 62.35 ppmv 1 - 31 Jul -- 1.57 μm Line Min: 3.17 ppmv \_\_\_\_ Max: 65.30 ppmv 20.0 17.5 15.0 2.05µm region 12.5 1 - 31 Jan -- 2.05 μm Line Min: 3.17 ppmv Max: 111.0 1 - 31 Jul -- 2.05 µm Line Error Max: 111.07 ppmv Min: 3.07 ppmv Max: 116.32 ppmv 10.0 has largest measurement 5.0 2.5 errors Measurement Error Ratio, 2.05/1.57 μm Min: 0.89 Max: 4.13 Measurement Error Ratio, 2.05/1.57 µm Min: 0.45 Max: 4.08 Max: 4.13 2.0 1.6 2.05/1.57 µm Raf 0.4 aer 0.0

Atmospheric and Environmental Research

© Atmospheric and Environmental Research, Inc. (AER), 2012.

# WRF-STILT

WRF [Skamarock and Klemp, 2008] :

•Mesoscale meteorological model provides transport fields

**STILT** [*Lin et al.,* 2003] **:** 

•A Lagrangian (airmass-following) transport model allowing backward in time transport simulations (receptor-oriented).

•Minimizes numerical diffusion present in Eulerian models [*Chevallier et al.,* 2007].

•Efficient way to calculate adjoint ("footprint") at high spatial and temporal resolution.

## WRF-STILT Coupling [Nehrkorn et al., 2010]:

•Realistic treatment of convective fluxes.

•Good mass conservation properties.



# **WRF-STILT** Application

## **Identify Receptor List**

- •Receptors represent ASCENDS observation locations
- •Receptor selection performed using CALIPSO orbital data over North America
  - •Optical depth < 0.7
  - •Surface detection frequency > 0
  - •~6,000 unique locations per day for Jan, Apr, Jul, and Oct 2007
  - •Expand receptor list to 15 vertical levels (0.5 14.5km)

## **Trajectory Simulations**

•WRF-STILT simulates release of 500 particles at each of the 15 levels for each receptor.

•Transport modeled backwards in time for 10 days.



# **WRF-STILT** Application

Description

## STILT •Revision 640 (<u>www.stilt-model.org</u>)

Ontion

## **WRF Physics and Numerics**

## WRF

•Version 2.2

•40km resolution

•NARR 32km grids

for IC and BC

| Option        | Description  |
|---------------|--|
| Land-surface  | Noah land-surface model with Monin-Obukov surface layer                          |
| PBL package   | Yongsei University (YSU) scheme  |
| LW radiation  | RRTM   |
| SW radiation  | Goddard  |
| Microphysics  | Lin et al.   |
| Convection    | Grell-Devenyi  |
| Nudging       | u,v,T,q at all levels above PBL, every 3 hours, 1 hour relaxation time           |
| Time stepping | 3 <sup>rd</sup> order Runge-Kutta  |
| Advection     | 5 <sup>th</sup> order horizontal, 3 <sup>rd</sup> order vertical                 |
|               | positive definite advection for moisture and scalars                             |
| Diffusion     | 2 <sup>nd</sup> order horizontal diffusion using Smagorinsky first-order closure |
| Damping       | No upper level or vertical velocity damping; default values for divergence and   |
|               | external model damping   |

•30 hr forecasts, reinitialized every 24 hrs with 6 hrs spin-up.

•Hourly output fields fed to STILT.



# Footprint

•Quantitatively describes how much of total mixing ratio at a receptor comes from surface fluxes originating in upwind regions

•Generated by WRF-STILT

•Units: mixing ratio per unit flux

$$\left[\frac{ppmv}{\mu mol \ / \ m^2 \ / \ s}\right]$$

•Domain: 11-65N, 50-170W

•Resolution: 1 deg horizontal & 3 hourly, 10 days back

•One unique location has 15 footprint maps – one for each vertical level

# **Vertically Integrated Footprint**

•Quantitatively describes how much of <u>column mixing ratio</u> comes from surface fluxes originating in upwind regions.

•Convolve 15 levels of footprint maps for each receptor with lidar weighting functions.



## **Vertically Integrated Footprints**



© Atmospheric and Environmental Research, Inc. (AER), 2012.

aer

**Environmental Research** 

# Formally Relate Footprint and Column CO<sub>2</sub>

$$\Delta C(x_r, y_r, t_r) = \sum_{v \in g}^{N} \prod_{n=1}^{15km} w(z_r) \iiint_{x \ y \ t} f(x, y, t | x_r, y_r, t_r, z_r) * \varpi_n(x, y) * [F_n^{prior} + d_n] * dxdydtdz$$

$$+ \int_{z_r}^{15km} w(z_r) \iiint_{x \ y \ t} f(x, y, t | x_r, y_r, t_r, z_r) * F_{anthro}^{prior} * \alpha_{anthro} * dxdydtdz_r$$
Simplifies to
$$\Delta C(x_r, y_r, t_r) = \sum_{v \in g}^{N} K_n(x_r, y_r, t_r) * [F_n^{prior} + d_n] + K_{anthro} * \alpha_{anthro}$$

GREEN = vegetative BLUE = anthropogenic  $\Delta C(x_r, y_r, t_r)$  =change in CO2 column mixing ratio at  $(x_r, y_r, t_r)$   $f(x, y, t | x_r, y_r, t_r, z_r)$  = footprint matrix at  $(x_r, y_r, t_r, z_r)$   $w(z_r)$  = 1 of 3 possible weighting functions  $\varpi_n(x, y)$  = vegetative coverage fraction at (x, y)

aer

Atmospheric and Environmental Research  $F_{anthro}^{prior}$  = a priori anthropogenic flux  $F_n^{prior}$  = prior flux for vegetation type n  $\alpha_{anthro}$  = multiplicative anthropogenic flux factor  $\alpha_n$  = additive flux factor for vegetation type n

# K<sub>anthro</sub> & K<sub>n</sub>

## **11 Vegetation Types Fraction Coverage**

## Anthropogenic (K<sub>anthro</sub>)

•Calculate using fluxes from the Vulcan Project anthropogenic emission inventory [*Gurney et al.,* 2009]

## Vegetative (K<sub>n</sub>)

•Calculate based on 11 vegetation types defined in in the Vegetation Photosynthesis and Respiration Model (VPRM) [*Mahadevan et al.,* 2008]





# A Posteriori CO<sub>2</sub> Flux Error

Apply integrated footprints  $K_n$  and  $K_{anthro}$  to calculate *a posteriori* CO<sub>2</sub> flux error due to introduction of ASCENDS observations:

$$\hat{S}_{post} = \left(\hat{K}^T \hat{S}_{\varepsilon}^{-1} \hat{K} + \hat{S}_{prior}^{-1}\right)^{-1}$$

$$\hat{S}_{post} = a \text{ posteriori flux errors } (\sigma_{post}^2)$$

$$\hat{K} = vegetative (K_n) \text{ or anthropogenic } (K_{anthro}) \text{ integrated footprint}$$

$$\hat{S}_{\varepsilon} = measurement \text{ error matrix}$$

$$\hat{S}_{prior} = a \text{ posteriori flux errors } (\sigma_{prior}^2)$$



# **A Priori Flux Errors**

## Vegetative

# •Apply $\hat{S}_{nrior}$ calculated by *Matross* [2006] to Jan & Jul 2007

- •Calculated for 11 vegetation types in VPRM
- •Off diagonal elements assumed 0
- •Diagonal elements estimated for summer 2004

## Anthropogenic

•Assume monthly mean *a priori* fractional error = 0.5 (overly pessimistic for Vulcan?)



## **Measurement Errors**

## •Best case scenario

•Consider contributions from random instrument measurement errors

•Exclude contributions from: atmospheric transport

lateral boundary condition

fossil fuel signal

[Gerbig et al., 2003]

## Calculation

•Diagonal matrix computed using reference value at Railroad Valley, NV (RRV)

 $\delta X_i$  = measurement error at receptor  $r_s(i)$  = surface reflectivity at recptor  $\tau_{atm}^2(i)$  = atmospheric transmission at recptor

$$\delta X_i = \delta X_0 \left[ \frac{r_s(i)}{r_s(0)} \frac{\tau_{atm}^2(i)}{\tau_{atm}^2(0)} \right]^{-1/2}$$

 $\delta X_0$  = measurement error at RRV  $r_s(0)$  = surface reflectivity at RRV  $\tau_{atm}^2(0)$  = atmospheric transmission at RRV

Surface reflectivity from MODIS backscatterOptical depth from CALIPSO data



## **Measurement Errors**



Atmospheric and **Environmental Research** 

aer

© Atmospheric and Environmental Research, Inc. (AER), 2012.

# **Baseline Results**

1.) Types with large *a priori* have large posterior error reduction on small timescale

2.) Types with small fractional coverage have smaller reductions

3.) 3pm offset doeslittle to improve upon1.57µm at 10 pm offset





# **Baseline Results**

4.) Greater flux error reductions for 2.05µm vs 1.57µm lines, particularly in January

5.) Modest increases in measurement error for 2.05µm line (blue) mask error reduction differences between lines

6) Flux error reductions are larger in July than January due to seasonality of measurement errors and regardless of weighting function





## **Results – Impact of Prior**

1.) Most "prior-less" results converge with prior constraint counterparts

2.) Exceptions are those types with smaller fractional coverage.



# Results – Increasing Measurement Error At RRV

- 1.) Much smaller posterior 0.5ppmv error reductions as measurement error increased
- 2) Larger posterior errorreduction differences at2.5ppmv

For larger measurement errors, up to ~1.5-3.5ppmv the advantage of a lowerpeaking weighting function is more important to constraining fluxes



# Results – Increasing Measurement Error At RRV

3.) As the measurement error increases, the prior constraint begins to dominate, particularly within the first 5 to 10 days, and impacting January more than July

ler

Atmospheric and Environmental Research



## **Results – Increasing Measurement Error at RRV**

### **Posterior Error Reduction After 1 Week**

4) As expected, measurement error of 0.5ppmv achieves greatest posterior error reduction

5) Nonlinear behavior exhibited by most types except shrub and grass





Environmental Research

# Conclusions

### Baseline

- 1.) Types with large *a priori* have large posterior error reduction on small timescale
- 2.) Types with small fractional coverage have smaller reductions
- 3.) 3pm offset does little to improve upon 1.57µm at 10 pm offset

4.) Greater flux error reductions for 2.05µm vs 1.57µm lines, particularly in January but modest increases in measurement error for 2.05µm line mask error reduction differences between lines

5.) Flux error reductions are larger in July than January due to seasonality of measurement errors and regardless of weighting function

### **Prior Constraint**

6.) The prior constraint has a limited impact on results at measurement error of 0.5 ppmv at RRV

### Measurement Error

7.) As measurement error increases, there is a nonlinear decrease in reduction of posterior error

8.)For larger measurement errors, up to ~1.5-3.5ppmv the advantage of a lower-peaking weighting function is more important to constraining fluxes

9.) As the measurement error increases, the prior constraint begins to dominate, particularly within the first 5 to 10 days, and impacting January more than July



## Future

•The footprint library provides valuable input to cost-benefit analyses of ground-based versus space-borne approaches to Measurement, Reporting, and Verification (MRV).

•The footprint library can be utilized in many other applications because footprint maps are not dependent on:

- •vertical weighting function or averaging kernel
- •measurement random errors
- •a priori flux errors
- •surface cover type and inversion approach
- •molecular species (i.e., footprints can be applied to gases other than  $CO_2$ ).

•The footprint library underlying this study has been computed on a daily basis for January, April, July, and October 2007 and is available to researchers with access to Pleiades (please contact jel@aer.com for more information).



## Acknowledgements

•This work has been funded by the NASA Atmospheric  $CO_2$  Observations from Space Program (grant NNX10AT87G).

•ASCENDS Measurement Requirement Definition team, especially

- •Randy Kawa
- •Ed Browell
- •Jim Abshire
- •Bob Menzies
- •Berrien Moore
- •Dan Matross

•NASA Ames HEC facility staff



## **Thank You**

#### References

Caron, J. and Y. Durand (2009), Operating wavelengths optimization for a spaceborne lidar measuring atmospheric CO<sub>2</sub>, *Appl. Opt. 48*, 5413-5422.

Chevallier, F., F.-M. Bre on, and P. J. Rayner (2007), Contribution of the Orbiting Carbon Observatory to the estimation of CO<sub>2</sub> sources and sinks: Theoretical study in a variational data assimilation framework, *J. Geophys. Res.*, *112*, D09307, doi:10.1029/2006JD007375.

Gerbig, C., J. C. Lin, S. C. Wofsy, B. C. Daube, A. E. Andrews, B. B. Stephens, P. S. Bakwin, and C. A. Grainger (2003), Towards constraining regional scale fluxes of CO<sub>2</sub> with atmospheric observations over a continent: 2. Analysis of COBRA data using a receptor-oriented framework, *J. Geophys. Res.*, *108* (D24), 4757, doi:10.1029/2003JD003770.

Gurney, K.R., D.L. Mendoza, Y. Zhou, M.L. Fischer, C.C. Miller, S. Geethakumar, and S. de la Rue du Can (2009), High Resolution Fossil Fuel Combustion CO2 Emission Fluxes for the United States, *Environ. Sci. & Technol. 43* (14), 5535-5541, doi:10.1021/es900806c.

Kawa, S. R., D. J. Erickson III, S. Pawson, and Z. Zhu (2004), Global CO<sub>2</sub> transport simulations using meteorological data from the NASA data assimilation system, *J. Geophys. Res.*, 109, D18312, doi:10.1029/2004JD004554.

Lin, J. C., C. Gerbig, S. C. Wofsy, B. C. Daube, A. E. Andrews, K. J. Davis, and C. A. Grainger (2003), A near-field tool for simulating the upstream influence of atmospheric observations: The Stochastic Time-Inverted Lagrangian Transport (STILT) model, *J. Geophys. Res.*, *108(D16)*, 4493, doi:10.1029/2002JD003161.

Mahadevan, P., S. C. Wofsy, D. M. Matross, X. Xiao, A. L. Dunn, J. C. Lin, C. Gerbig, J. W. Munger, V. Y. Chow, and E. W. Gottlieb (2008), A satellite-based biosphere parameterization for net ecosystem CO<sub>2</sub> exchange: Vegetation Photosynthesis and Respiration Model (VPRM), *Global Biogeochem. Cycles*, *22*, GB2005, doi:10.1029/2006GB002735.

Matross, D. M. (2006), Regional Scale Land-Atmosphere Carbon-Dioxide Exchange: Data Design and Inversion within a Receptor Oriented Modeling Framework, PhD Thesis, Department of Earth and Planetary Sciences, Harvard University, 171 pp.

Nehrkorn, T., J. Eluszkiewicz, S. C. Wofsy, J. C. Lin, C. Gerbig, M. Longo, and S. Freitas (2010), Coupled Weather Research and Forecasting/Stochastic Time-Inverted Lagrangian Transport (WRF-STILT) model, *Meteor. Atmos. Phys., 107 (1-2),* 51–64, doi:10.1007/s00703-010-0068-x.

Skamarock, W. C. and J. B. Klemp (2008), A time-split nonhydrostatic atmospheric model for weather research and forecasting applications, *J. Comp. Phys.*, 227, 3465-3485.



## **Backup Slides**





## **Backup Slides**



